

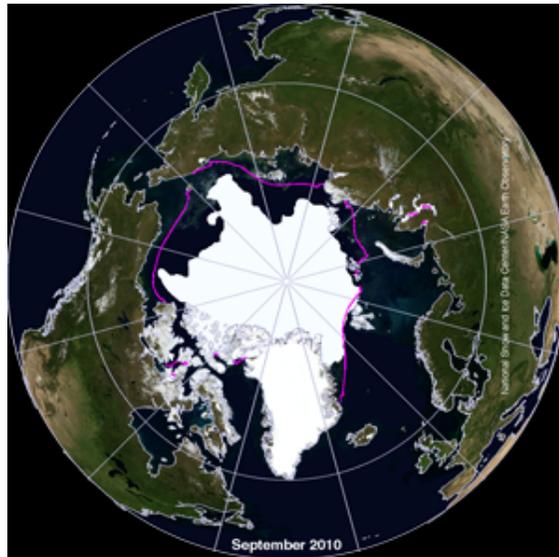
Outline

- 1. Introduction**
- 2. Theoretical background**
- 3. Retrieval method**
- 4. Data and procedure**
- 5. Results**
- 6. Summary and conclusion**





- Sea ice is one of the most important variables in the global climate system, and in understanding the surface energy budget, fresh-water, and global temperature.
- Sea ice melting impacts on Ocean circulation, Melting of the Greenland Ice Sheet, Changes of animals, in vegetation, sea floor methane(Greenhouse gas), territorial claims (resources development, new shipping routes, sovereign rights)



<<http://nsidc.org/icelights/crash-course/arctic-sea-ice/>>



<earthobservatory.nasa.gov>

- ◆ Sea ice extent was a minimum in Sept. 2007.
- ◆ This research is in review.



- Arctic shortcut in the Northern Sea
 - China (Aug. 21 2013)
 - : Sailing a container-transporting vessel along the Northern Sea Route
 - Korea and Japan
 - : Shipping oil products through the Arctic's melting ice along the Northern Sea Route.

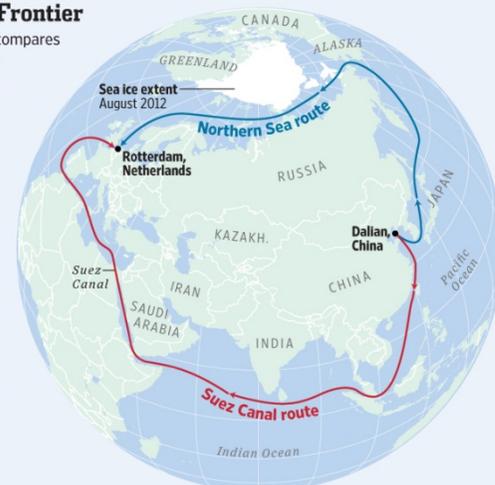
- Arctic council

- Member states
 - : Canada, Finland, Iceland, Norway, Russia, Sweden, United States, Denmark; representing also the dependencies of Greenland and Faroe Islands
- Permanent observer states
 - : China, France, Germany, India, Italy, Japan, South Korea, Netherlands, Poland, Singapore, Spain, United Kingdom

China's New Shipping Frontier

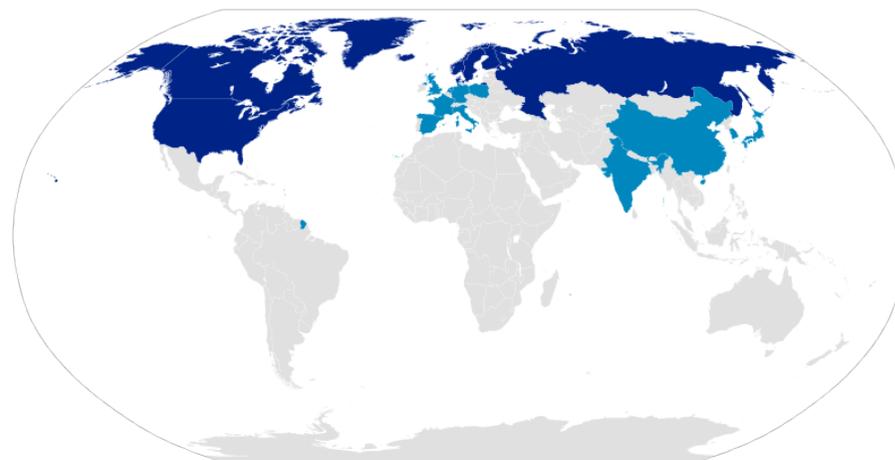
How the new Northern Sea route compares to the traditional Suez Canal route

NORTHERN SEA ROUTE	
Travel time	35 days
Dangers	Icebergs
Travel window	July to November
Container-carrying vessels	One this year
SUEZ CANAL ROUTE	
Travel time	48 days
Dangers	Access to Suez Canal under question with upheaval in Egypt
Travel window	Year-round
Container-carrying vessels	17,000 last year



Sources: Northern Sea Route Information Office; National Snow and Ice Data Center; Cosco; Lloyd's List The Wall Street Journal

<http://online.wsj.com/article/SB10001424127887324619504579026031203525734.htm?project%3DARCTICSHIP%26articleTabs%3Dinteractive>





- Microwave satellite remote sensing is useful to monitor the change of sea ice because a large contrast in emissivity (or reflectivity) exists between sea ice and water.



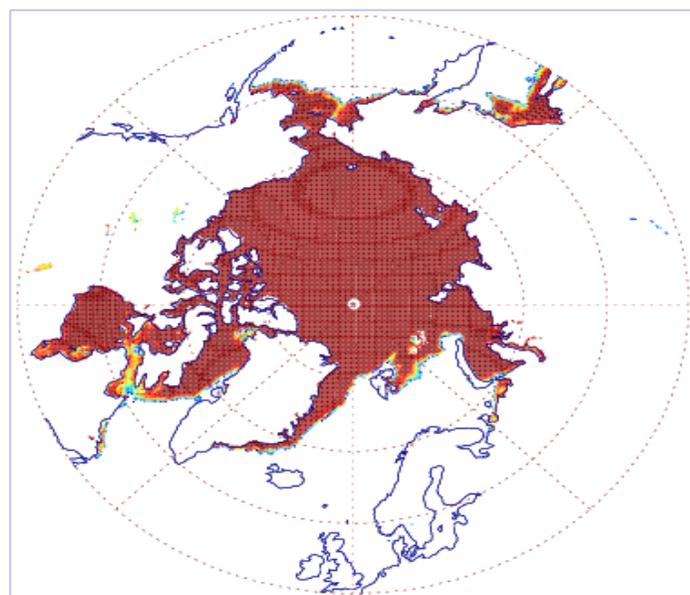
AQUA/AMSR-E sensor♪

Table 2.1 – AMSR-E performance characteristics

CENTER FREQUENCIES (GHz)	6.925	10.65	18.7	23.8	36.5	89.0
BANDWIDTH (MHz)	350	100	200	400	1000	3000
SENSITIVITY (K)	0.3	0.6	0.6	0.6	0.6	1.1
MEAN SPATIAL RESOLUTION (km)	56	38	21	24	12	5.4
IFOV (km x km)	74 x 43	51 x 30	27 x 16	31 x 18	14 x 8	6 x 4
SAMPLING RATE (km x km)	10 x 10	5 x 5				
INTEGRATION TIME (MSEC)	2.6	2.6	2.6	2.6	2.6	1.3
MAIN BEAM EFFICIENCY (%)	95.3	95.0	96.3	96.4	95.3	96.0
BEAMWIDTH (degrees)	2.2	1.4	0.8	0.9	0.4	0.18



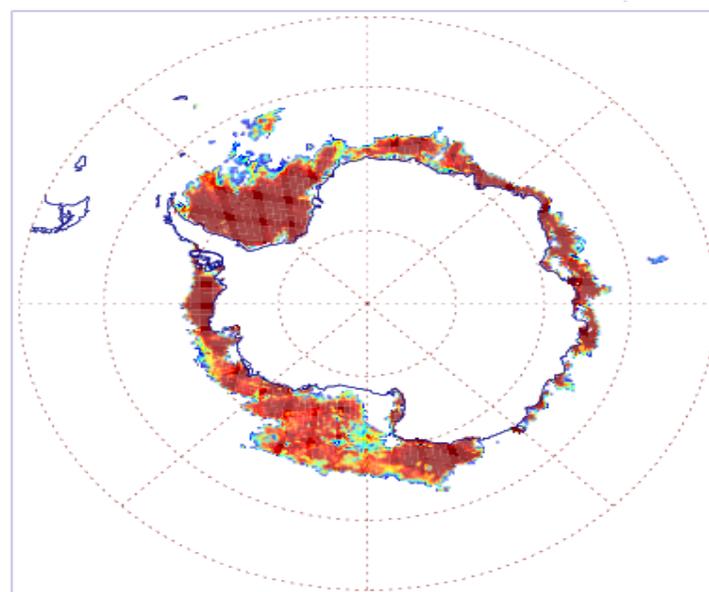
- Sea ice concentration is used in characterizing the spatial sea ice cover and extent and in the long-term trend analyses and process studies (Bjørge et al. 1997; Cavalieri et al. 1997; Parkinson et al. 1999; Zwally et al. 2002)



Sea ice concentration (%)

15.0 32.0 49.0 66.0 83.0 100.0

<Jan. 1 2007, Arctic region >



Sea ice concentration (%)

15.0 32.0 49.0 66.0 83.0 100.0

<Jan. 1 2007, Antarctic region>



- Satellite microwave radiometer measure the energy emitted from the ocean surface.
 - The intensity of polarized radiation can be observed by microwave sensor depends on the **sea surface emissivity (or reflectivity)** .

- In this research, we investigate the following surface properties of sea ice:
 1. Small-scale surface roughness
 2. Refractive index of sea ice

- We propose an alternative approach for retrieving the sea ice properties using
 1. Hong approximation (Hong, 2009, J. Appl. Rem. Sens.)
 2. ASH(Azzam-Sohn-Hong) approximation (Hong, 2013, IEEE Geosci. Rem. Sens. Lett.)

The objective of this study is to improve retrieval of sea ice surface properties from satellite passive-microwave observations



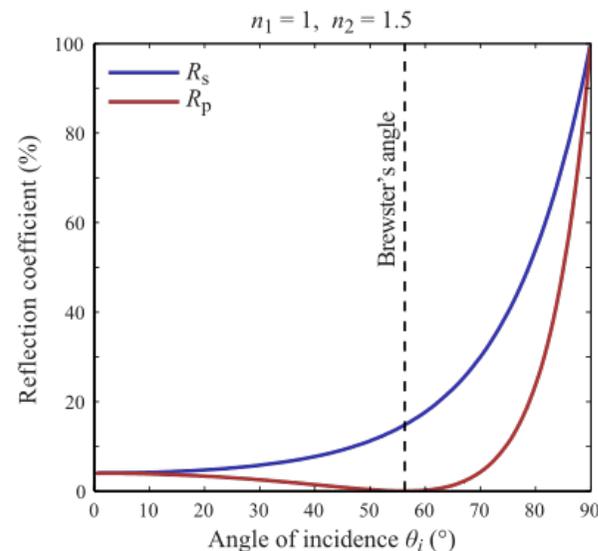
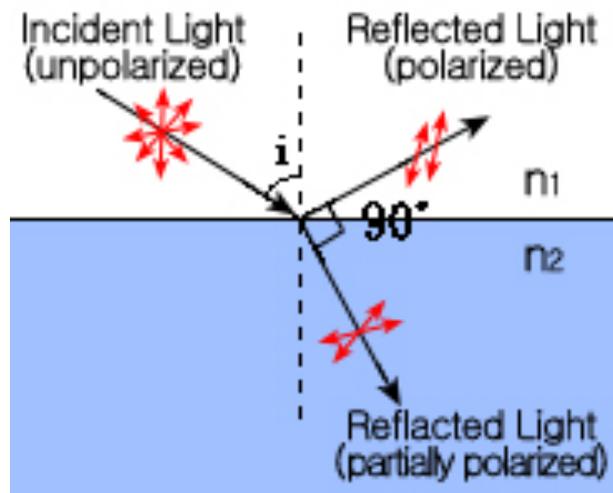
■ Fresnel equation

$$R_v = \left| \frac{\cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2 \quad \text{and} \quad R_h = \left| \frac{\hat{n}^2 \cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\hat{n}^2 \cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2$$

where θ is the view angle, $\hat{n} = n + ik$ ($i = \sqrt{-1}$) is the complex refractive index of the material and n and k are the real and imaginary parts of the refractive indices, respectively.



Augustin Fresnel
(1788-1827)





- **Questions**

(1) Is there any relationship between polarized reflectivities if the refractive index is unknown?

(2) How can we estimate the refractive index of a material?

*** We measure the reflectivities in a real world**

*** Refractive index is a complex number**



- **Fresnel equation**

$$R_v = \left| \frac{\cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2 \quad \text{and} \quad R_h = \left| \frac{\hat{n}^2 \cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\hat{n}^2 \cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2$$

- Constraints:

(1) **Two non-linear equations (← Fresnel equation)**

(2) **Three unknown variables (angle, real and imaginary parts of the refractive index) $\tilde{n} = n + i\kappa$.**

➔ **Difficult problem to be solved !**

➔ **Therefore, I tried to find a relationship between two polarizations irrespective of refractive index**



- **Hong approximation** (Hong, 2009)
- **The generalized Fresnel equation** (Tousey, 1935)

$$R_V = \frac{(A_+ - \cos \theta)^2 + A_-^2}{(A_+ + \cos \theta)^2 + A_-^2} \quad \text{and} \quad R_H = R_V \times \frac{(A_+ - \sin \theta \tan \theta)^2 + A_-^2}{(A_+ + \sin \theta \tan \theta)^2 + A_-^2},$$

with $A_{\pm}^2 = [\sqrt{(n^2 - K^2 - \sin^2 \theta)^2 + 4n^2 k^2} \pm (n^2 - K^2 - \sin^2 \theta)] / 2,$

- From taking the first term in the Taylor series for natural logarithmic ratio of $\ln R_H / \ln R_V$, we can obtain the Hong approximation as follows (Hong, 2009):

$$\frac{\ln R_H}{\ln R_V} = 1 + \frac{\ln(1 + X_3) - \ln(1 + X_4)}{\ln(1 + X_1) - \ln(1 + X_2)}$$

with $X_1 = (Q - \cos \theta)^2 + P^2 - 1,$ $X_2 = (Q + \cos \theta)^2 + P^2 - 1,$

$X_3 = (Q - \sin \theta \tan \theta)^2 + P^2 - 1,$ and $X_4 = (Q + \sin \theta \tan \theta)^2 + P^2 - 1.$

After taking the first term in the Taylor series for natural logarithm ($\ln(1+x) = x - x^2/2 + x^3/3 + O(x^4)$) without the strict constraints ($-1 < x \leq 1$), the approximation of Eq. (2) is as follows:

$$\frac{\ln R_H}{\ln R_V} \approx 1 + \frac{X_3 - X_4}{X_1 - X_2} = 1 + \frac{\sin \theta \cdot \tan \theta}{\cos \theta} = \frac{1}{\cos^2 \theta}$$

$$R_H = R_V^{1/\cos^2 \theta}$$



- **Hong approximation (Hong, 2009)**

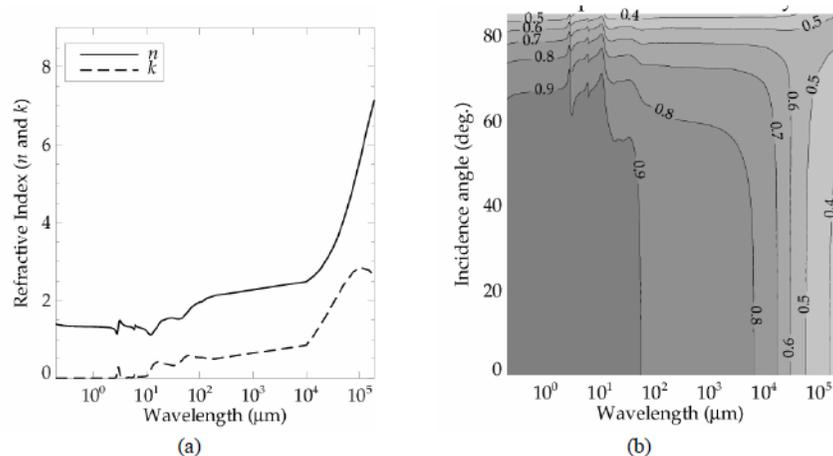


Fig. 1. (a) Refractive index of water and (b) unpolarized emissivity. The same wavelength range is used as that shown in Figure 1. The unpolarized emissivity is calculated using the Fresnel equations and the refractive index of water for various incidence angles.

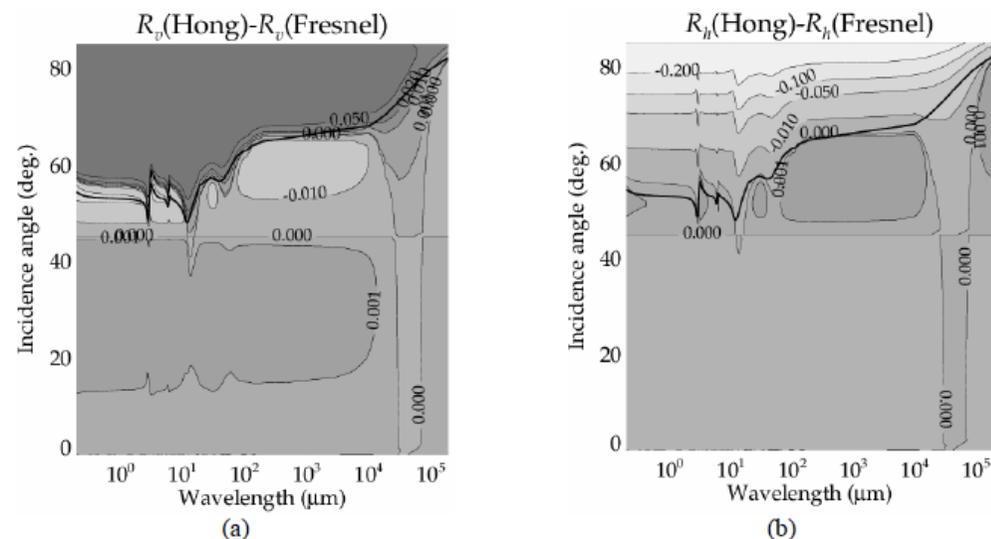


Fig. 2. Validation of the Hong approximation. (a) $R_V(\text{Fresnel}) - R_V(\text{Hong})$ and (b) $R_H(\text{Fresnel}) - R_H(\text{Hong})$ are shown. The wavelength range is the same as that shown in Fig. 1.



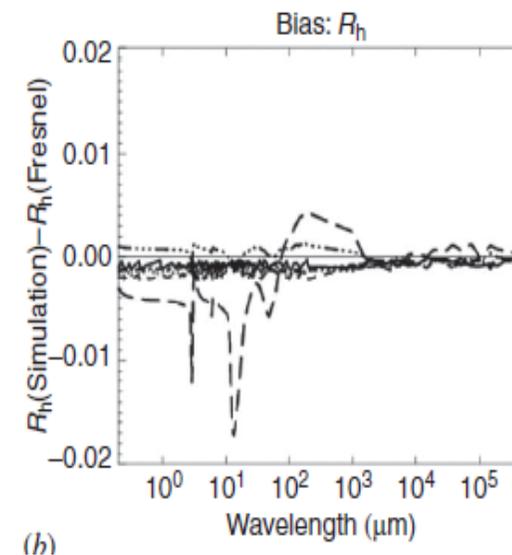
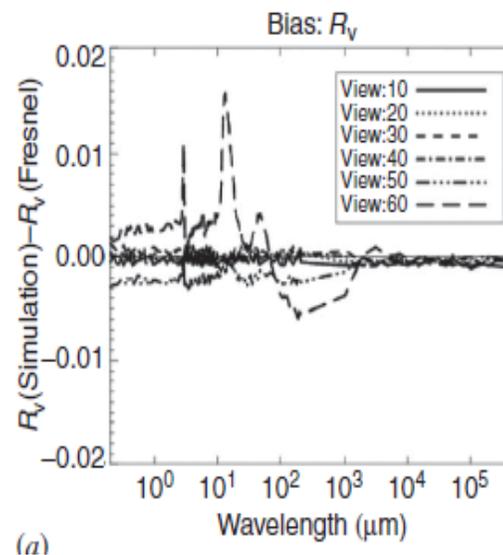
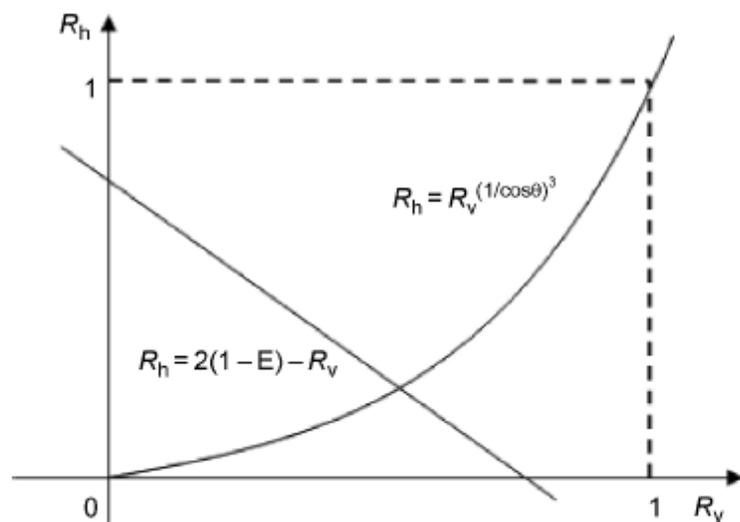
1. Decomposition of unpolarization

(1) Hong approximation

$$R_H = R_V^{1/\cos^2 \theta}$$

(2) Unpolarized reflectivity

$$R = \frac{R_V + R_H}{2},$$



Hong,2010 (Int. J. Rem. Sens.)



2. Refractive Index

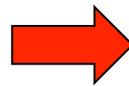
- : Derived a theoretical relationship between polarized reflectivities
- : Provide how to estimate the refractive index (intrinsic property of materials)

$$R_v = \left| \frac{\cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2 \quad \text{and} \quad R_h = \left| \frac{\hat{n}^2 \cos \theta - \sqrt{\hat{n}^2 - \sin^2 \theta}}{\hat{n}^2 \cos \theta + \sqrt{\hat{n}^2 - \sin^2 \theta}} \right|^2,$$

Query (1969)

$$n = \sqrt{\frac{A_+^2 - A_-^2 + \sin^2 \theta + \sqrt{(A_+^2 + A_-^2 - \sin^2 \theta)^2 + 4A_+^2 A_-^2}}{2}} \quad \text{and} \quad K = \frac{A_+ A_-}{n}, \quad (9)$$

where $A_+ = \sqrt{-A_+^2 - 2FA_+ \cos \theta - \cos^2 \theta}$, $A_- = [(F-G)\sin \theta \cot 2\theta] / [GF + (1-F^2)\cos^2 \theta - 1]$,
 $F = (R_v + 1) / (R_v - 1)$, and $G = (R_h + 1) / (R_h - 1)$.



(Hong, 2009)

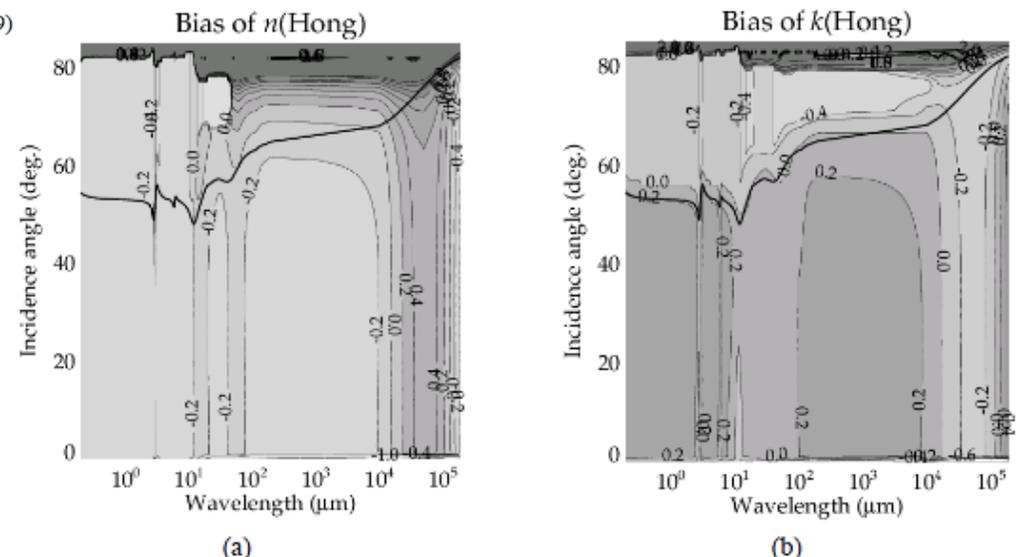


Fig. 4. Bias of the retrieved refractive index of water for (a) n and (b) K .



3. Surface Roughness

$$R_{R,V} \approx R_{S,V} \approx 0 \text{ and } R_{R,H} < R_{S,H} \quad R_{S,H} = R_{S,V}^{\cos^2 \theta}$$



$$\sigma = \frac{\lambda}{4\pi \cos \theta} \cdot \sqrt{\ln \left(\frac{R_{R,V}^{\cos^2 \theta}}{R_{R,H}} \right)}$$

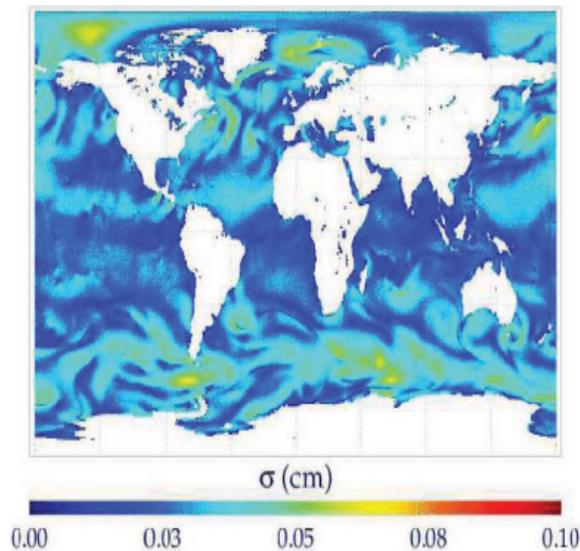


Figure 4. Sea surface roughness at SSM/I 19 GHz channel for 24 February 2009 (06 UTC). FASTEM-2, Fresnel equations, equation (1), and GDAPS data of 24 February 2009 were used. The small-scale surface roughness ranged from 0 to 0.07 cm in this case.

Hong, 2010 (Int. J. Rem. Sens.)

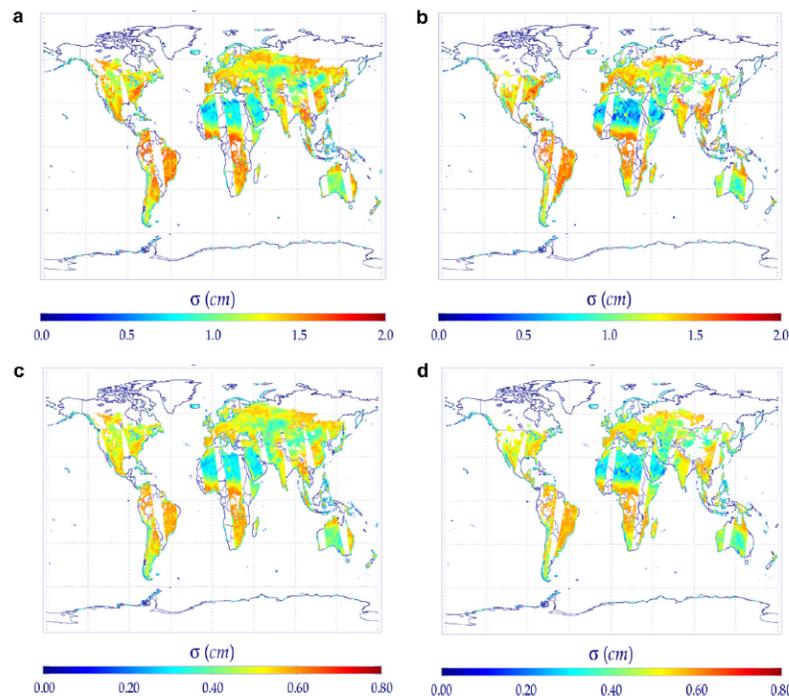
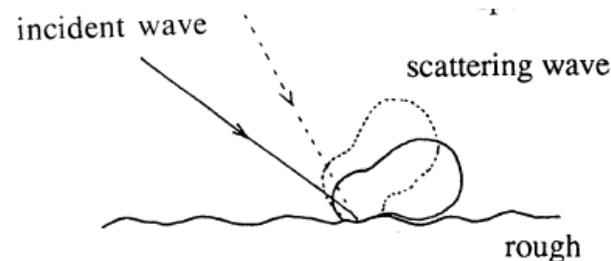
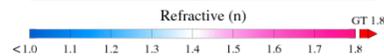
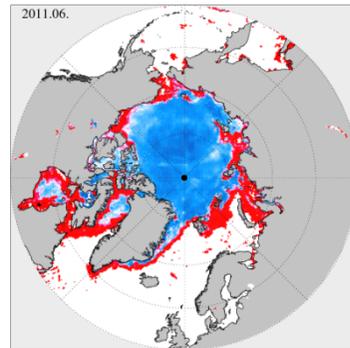
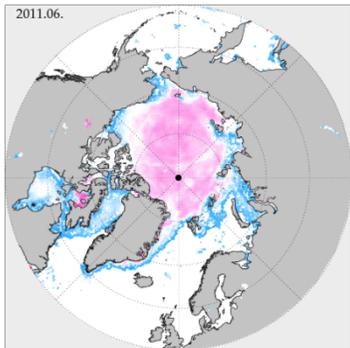
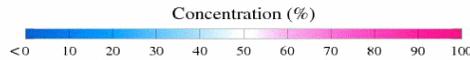
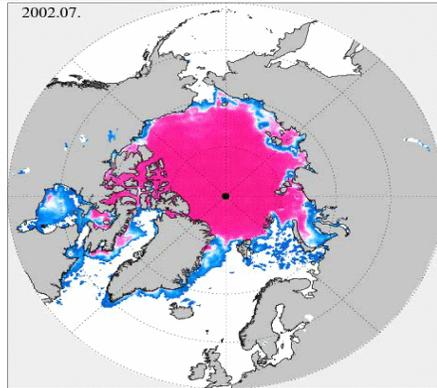


Fig. 5. Estimated σ using W-C model at AMSR-E (a) 6.9 GHz (ascending), (b) 6.9 GHz (descending), (c) 10 GHz (ascending), and (d) 10 GHz (descending). The σ is relatively low for the desert and bare soil regions, and relatively high for vegetation regions.

Hong, 2010 (J. Hydrol.)

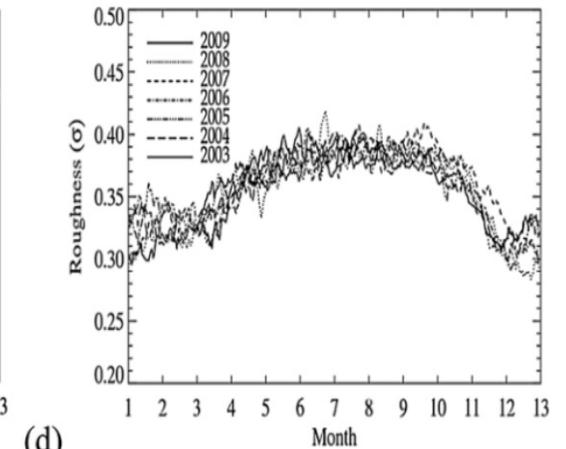
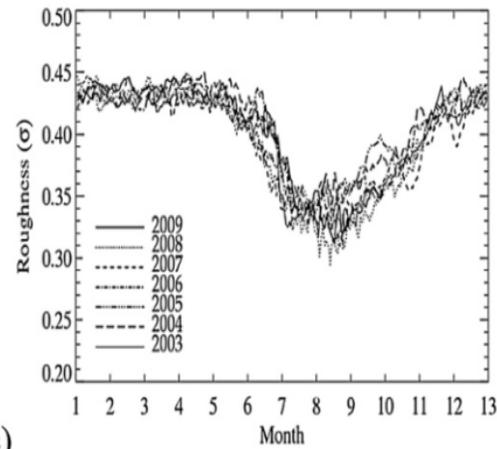
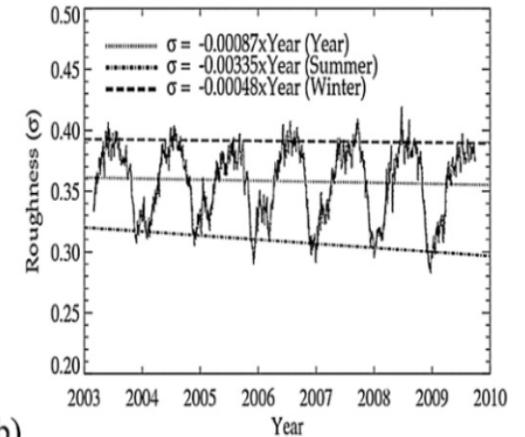
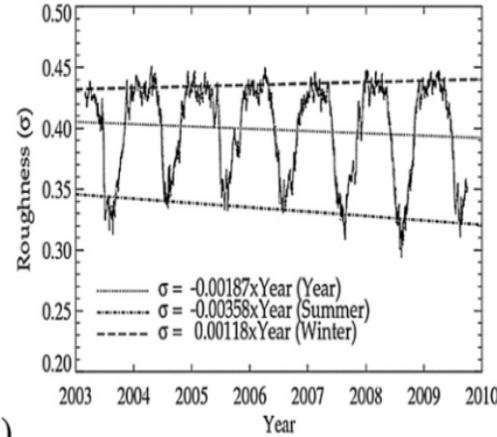


4. Sea ice



Concentration, Small-scale roughness and refractive index of sea ice surfaces (June, 2011) using AQUA/AMS R-E data

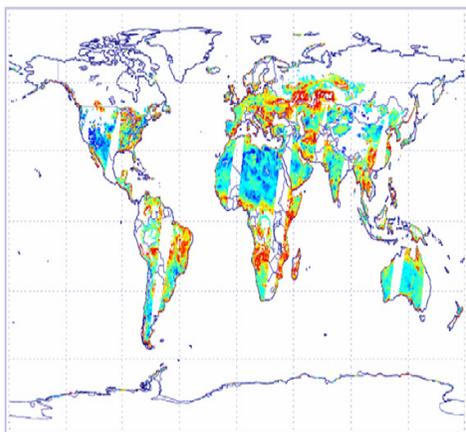
Hong, 2010 (Rem. Sens. Environ.)



Hong and Shin, 2010 (J. Climate)

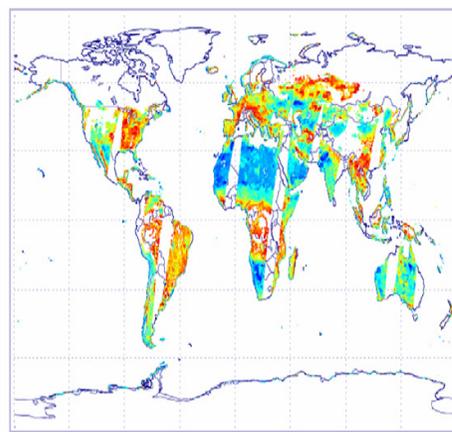


5. Soil moisture



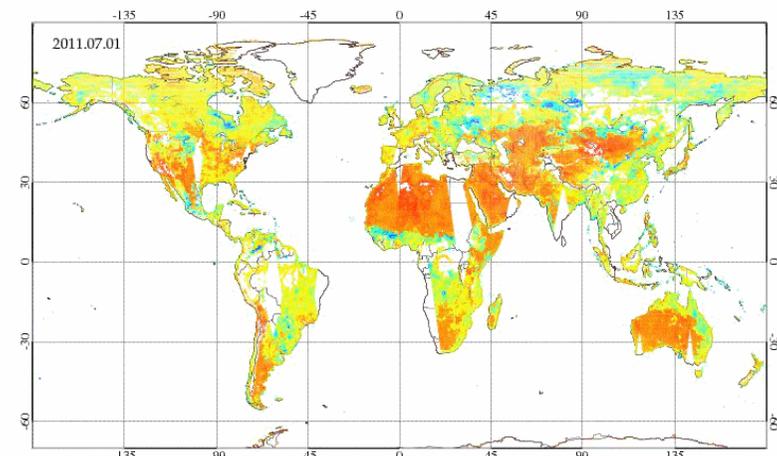
Volumetric soil water content (m^3/m^3)

0.00 0.04 0.08 0.12 0.16 0.20



Volumetric soil water content (m^3/m^3)

0.00 0.04 0.08 0.12 0.16 0.20



Volumetric soil water content (m^3/m^3)

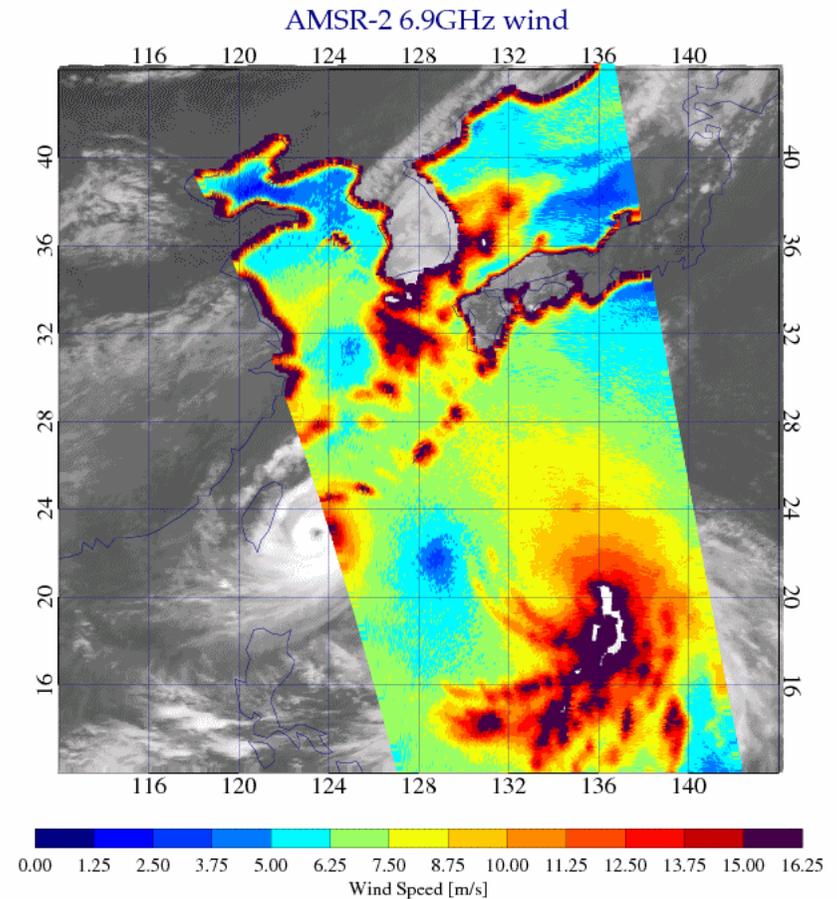
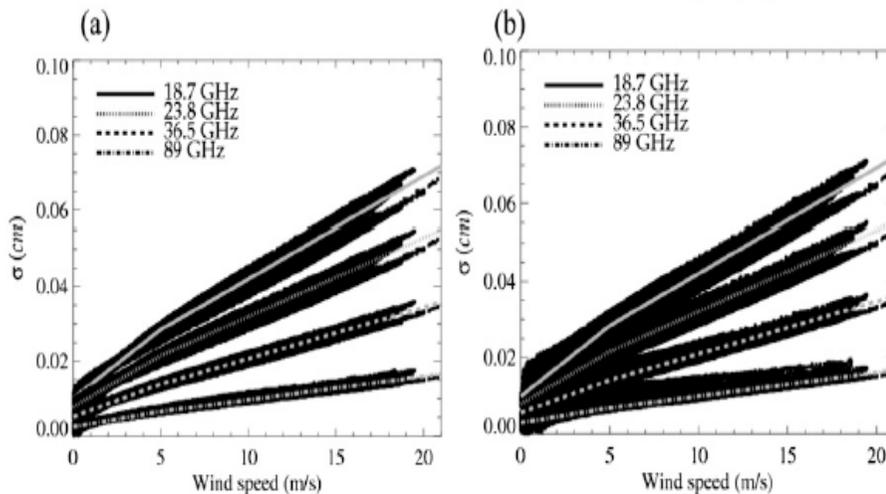
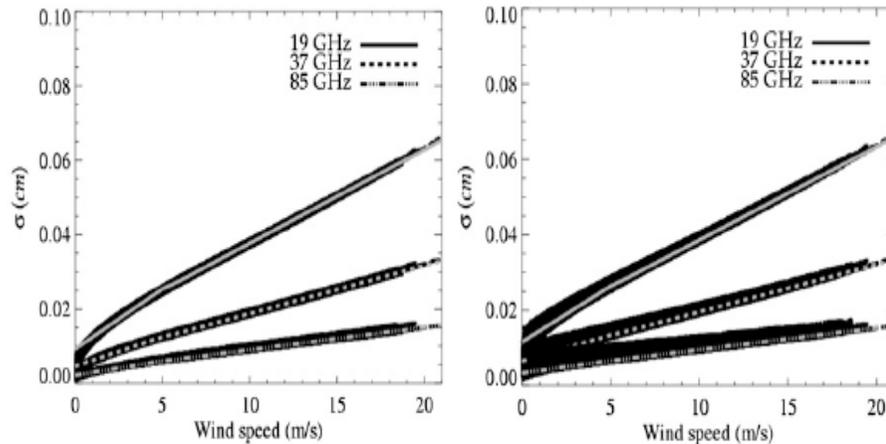
0.00 0.07 0.14 0.21 0.28 0.35

- (a) Volumetric soil water content (g/cm^3) using Hong's algorithm
- (b) AMSR-E level 3 products.

Hong and Shin, 2011 (J. Hydrol.)



6. Wind speed



- Hong and Shin (2013)



- **Direct relationship (Azzam relationship) (Hong, 2013 (IEEE GRSL))**

$$R_{S,H} = R_{S,V} \times \frac{(\sqrt{R_{S,V}} - \cos 2\theta)^2 + 2(\sqrt{R_{S,V}} - \text{Re}(\hat{r}_V) \cdot \cos 2\theta)}{(1 - \sqrt{R_{S,V}} \cdot \cos 2\theta)^2 + 2(\sqrt{R_{S,V}} - \text{Re}(\hat{r}_V) \cdot \cos 2\theta)}$$

where $R_{S,V} = \text{Re}(\hat{r}_V)^2 + \text{Im}(\hat{r}_V)$ such that $\text{Re}(\hat{r}_V)$ and $\text{Im}(\hat{r}_V)$ are the real and imaginary parts of \hat{r}_V , respectively.



Azzam (1979, 1986)

Starting from (1) by substituting \hat{n}^2 into (2) and using $\cos^2 \theta = (1 + \cos 2\theta)/2$, the following relationship between the complex Fresnel reflection coefficients has been obtained [27], [28]

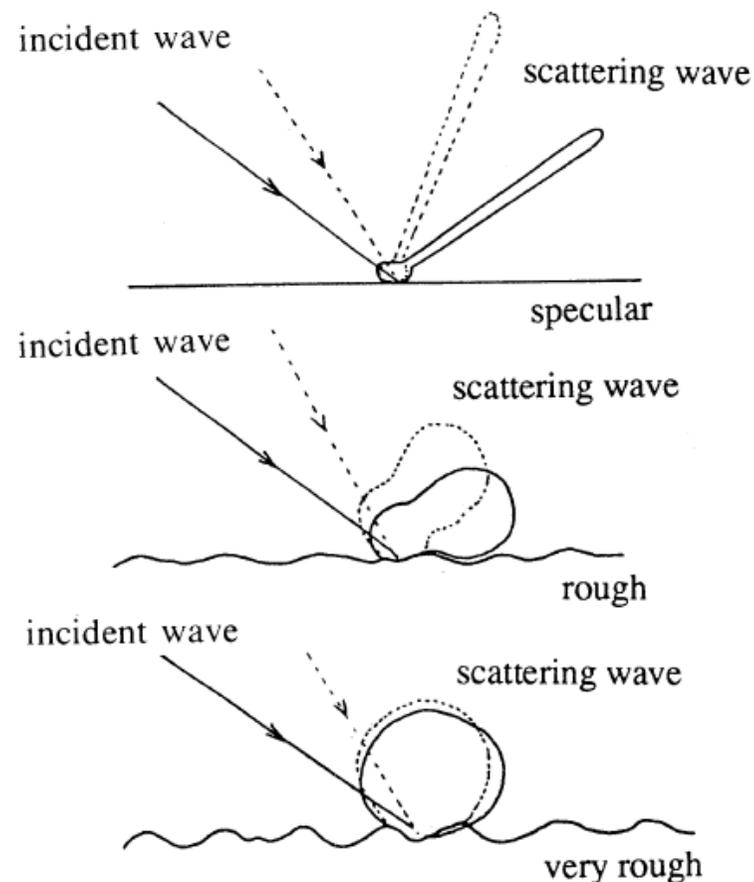
$$\hat{r}_H = \hat{r}_V \cdot \frac{\hat{r}_V - \cos 2\theta}{1 - \hat{r}_V \cdot \cos 2\theta}$$

- Weakness: it still needs *a priori* information
- In case of $k \sim 0$ (for example, dielectrics; For metals, $k > n$), Azzam relationship can be reduced to ASH (Azzam-Sohn-Hong) approximation as follows (Hong, 2013 (IEEE GRSL)):

$$R_{S,H} = R_{S,V} \cdot \frac{(\sqrt{R_{S,V}} + \cos 2\theta)^2}{(1 + \sqrt{R_{S,V}} \cdot \cos 2\theta)^2}$$



- **Roughness effect on the surface emissivity**
 - Passive microwave emission is affected by ice surface roughness as a complex function of the frequency, polarization, sensor-surface geometry of the radiometer, and complex permittivity of the material bounded below by the rough surface (Kong et al., 1979; Grenfell and Lohanick, 1985; Grenfell, 1992; Barber et al. 1998; Grody, 2008)



- **Scale vs. Wavelength**
 - Small-scale (roughness < wavelength)
 - Large-scale (roughness > wavelength)



- **Roughness effect on the surface emissivity**

- A semiempirical model based on the incoherent approach is used for small-scale surface roughness, which corresponds to identifying the height probability density function with a Gaussian distribution of zero mean and variance as follows (Wu & Fung, 1972, Choudhury et al., 1979):

$$R_R = R_S \cdot \exp\left[-\left(\frac{4\pi\sigma \cos\theta}{\lambda}\right)^2\right]$$

where R is the reflectivity, σ is the small-scale roughness height (Ulaby et al., 1982), λ is the wavelength, θ is the incidence angle. Subscripts R and S mean 'rough' and 'specular', respectively.

- **Small scale roughness retrieval**

- Hong approximation (Hong, 2009): $R_{S,H} \approx R_{R,V} \tan^2 2\theta$
- ASH approximation (Hong, 2103): $R_{S,H} = R_{S,V} \cdot \frac{(\sqrt{R_{S,V}} + \cos 2\theta)^2}{(1 + \sqrt{R_{S,V}} \cdot \cos 2\theta)^2}$
- Estimate using the observed rough surface reflectivities near the Brewster angle (Hong, 2010)

)

$$\text{Small scale Roughness: } \sigma \approx \lambda / 4\pi \cos\theta \cdot \sqrt{\ln(R_{R,V} \tan^2 2\theta / R_{R,H})}$$



- **Surface refractive index**

- Conversion rough surface reflectivity to specular surface reflectivity
- AMSR-E TB & Ts

→ Rough surface reflectivity

→ Specular surface reflectivity

→ Inversion method to estimate surface refractive index



Surface roughness

$$\sigma \approx \frac{\lambda}{4\pi \cos \theta} \cdot \sqrt{\ln \left(\frac{R_{R,v}^{\sec^2 \theta}}{R_{R,H}} \right)}$$

- **Retrieval of Surface refractive index** (Querry, 1969 (J. Opt. Soc. Am.))

$$n = \sqrt{\frac{A_+^2 - A_-^2 + \sin^2 \theta + \sqrt{(A_+^2 + A_-^2 - \sin^2 \theta)^2 + 4A_-^2 A_+^2}}{2}} \quad \text{and} \quad K = \frac{A_- A_+}{n},$$

$$\text{where } A_- = \sqrt{-A_+^2 - 2FA_+ \cos \theta - \cos^2 \theta},$$

$$A_+ = [(F - G) \sin \theta \cot 2\theta] / [GF + (1 - F^2) \cos^2 \theta - 1],$$

$$F = (R_V + 1) / (R_V - 1),$$

and

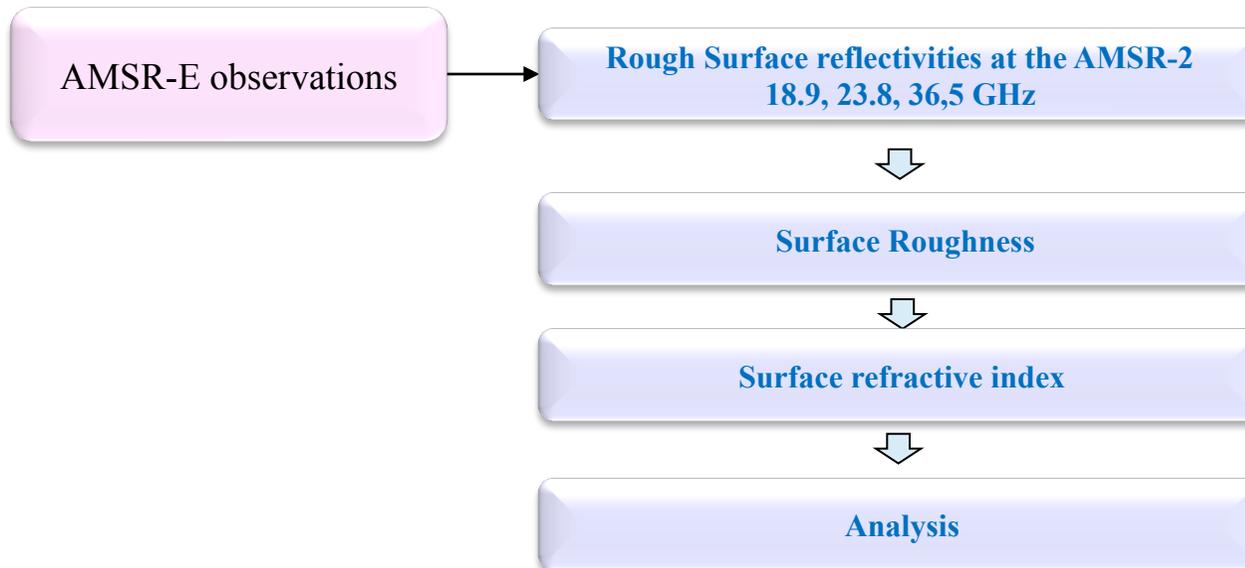
$$G = (R_H + 1) / (R_H - 1).$$



▪ Data

- AMSR-E level-3 data (25 km gridded data)
- channels : 6.925, 10.65, 18.7, 23.8, 36.5, and 89.0 GHz
- Jun. 2002 to Sep. 2011

▪ Procedure





- **Surface temperature (Comiso et al. 2003)**

$$T_S = T_{B,V}(6.9) / [1 - R_{R,V}(6.9)]$$

where $R_{R,P}(6.9)$ is 0.02 for ice (Markus and Cavalieri, 2000) and 0.44 for ocean (Hwang and Barber, 2008)

At other frequencies such as 18.7 GHz, the $R_{R,P}$ of the sea ice surface is estimated for polarization ($P = V$ or H) as follows:

$$R_{R,P}(18.7) = 1 - \frac{T_{B,P}(6.9)}{T_{B,P}(18.7)} [1 - R_{R,P}(6.9)]$$

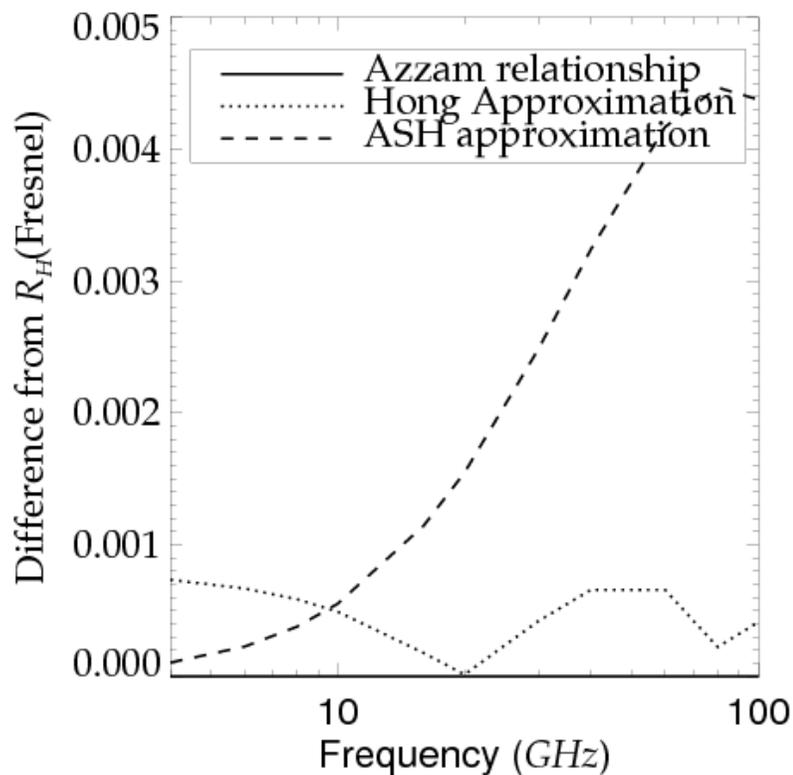
Finally, the mixed surface reflectivity of the ice and water on the rough sea ice surface is described as follows (Nishio and Comiso, 2005):

$$R_{R,V} = R_{R,V,ice} + (R_{R,V,ice} - R_{R,V,water}) \cdot \left[1 - \frac{C_I}{100} \right]$$

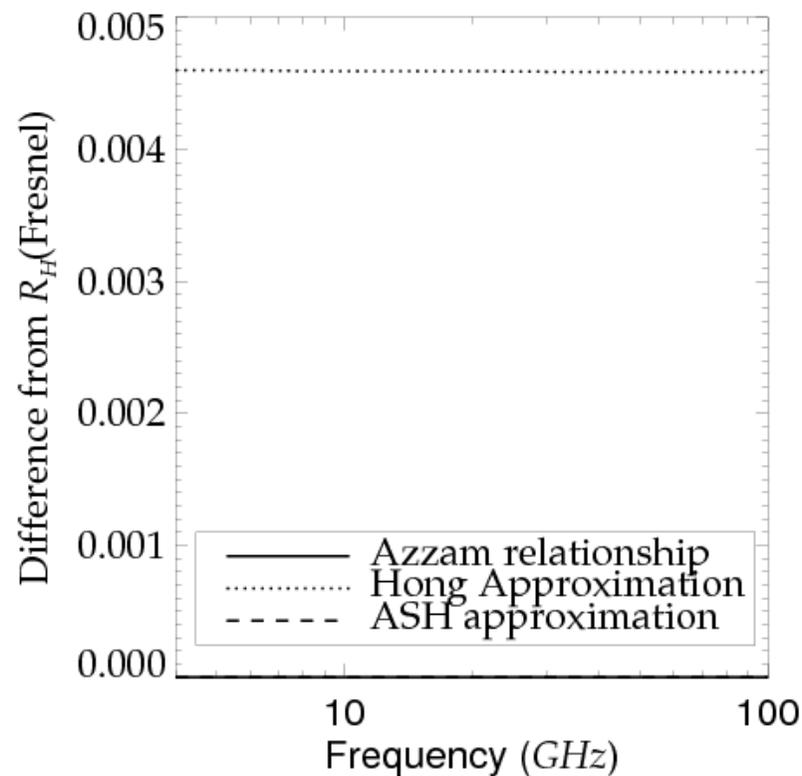
where C_I is defined as the percentage of sea ice within the field of view of the sensor (Nishio and Comiso, 2005).



- Absolute difference in R_H between the Fresnel equation and the ASH and Hong approximations for (a) ice and (b) water. AMSR-E has an incidence angle of 55° .



For water♪

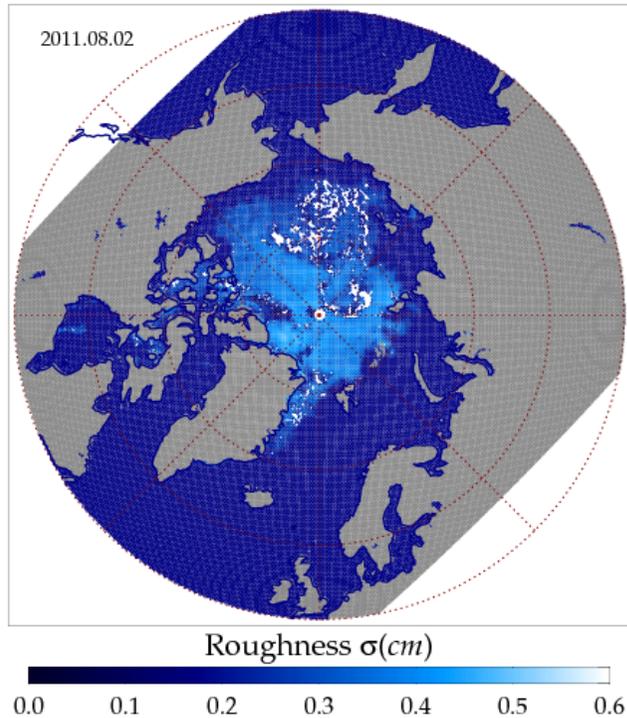


For ice♪

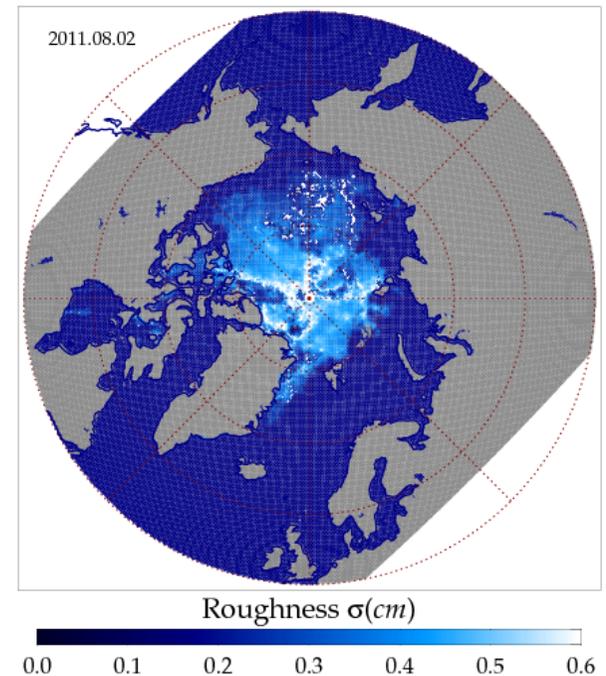
Results (roughness)



- Small-scale surface roughness σ retrieved for the Arctic using the Hong approximation and the ASH approximations in the 18.7 GHz channel of AMSR-E.



Hong approximation♪

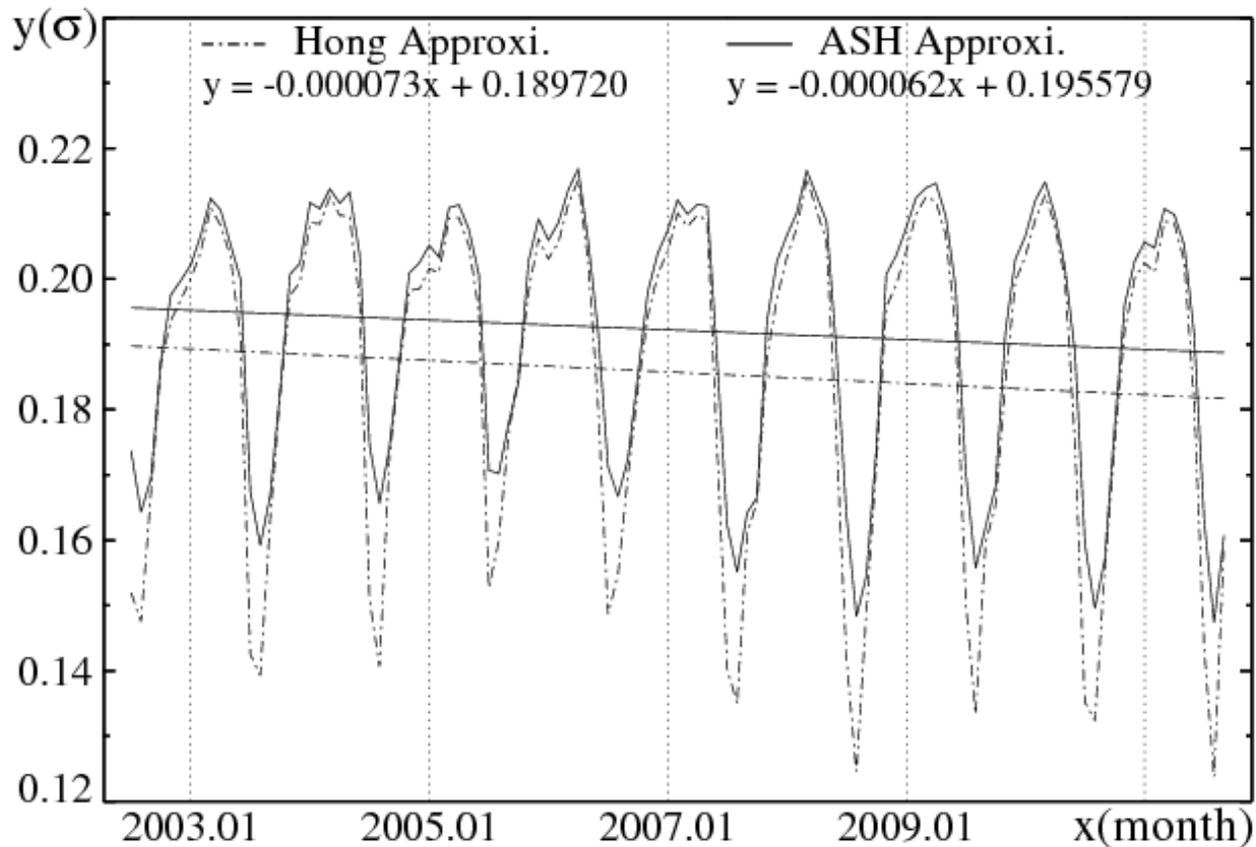


ASH approximation♪

Results (roughness)



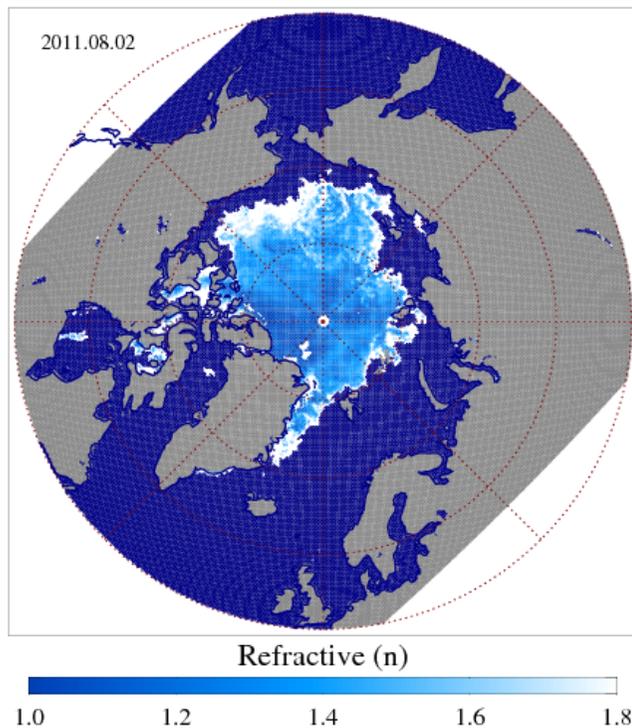
- The time series of σ from 2002 to 2011 show (c) the decreasing trend. The differences are relatively large in the melting seasons



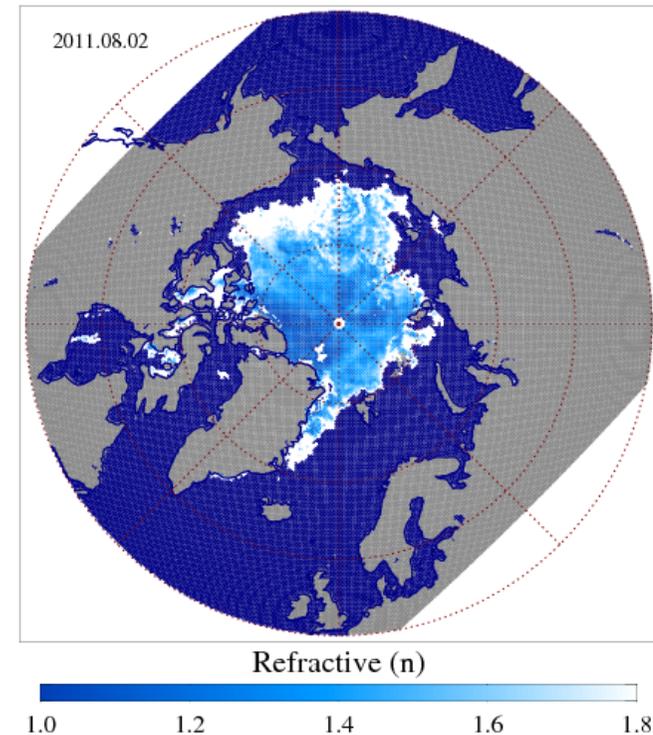
Results (Refractive index)



- Refractive index n retrieved for the Arctic using the Hong approximation and the ASH approximations in the 18.7 GHz channel of AMSR-E.



Hong approximation♪

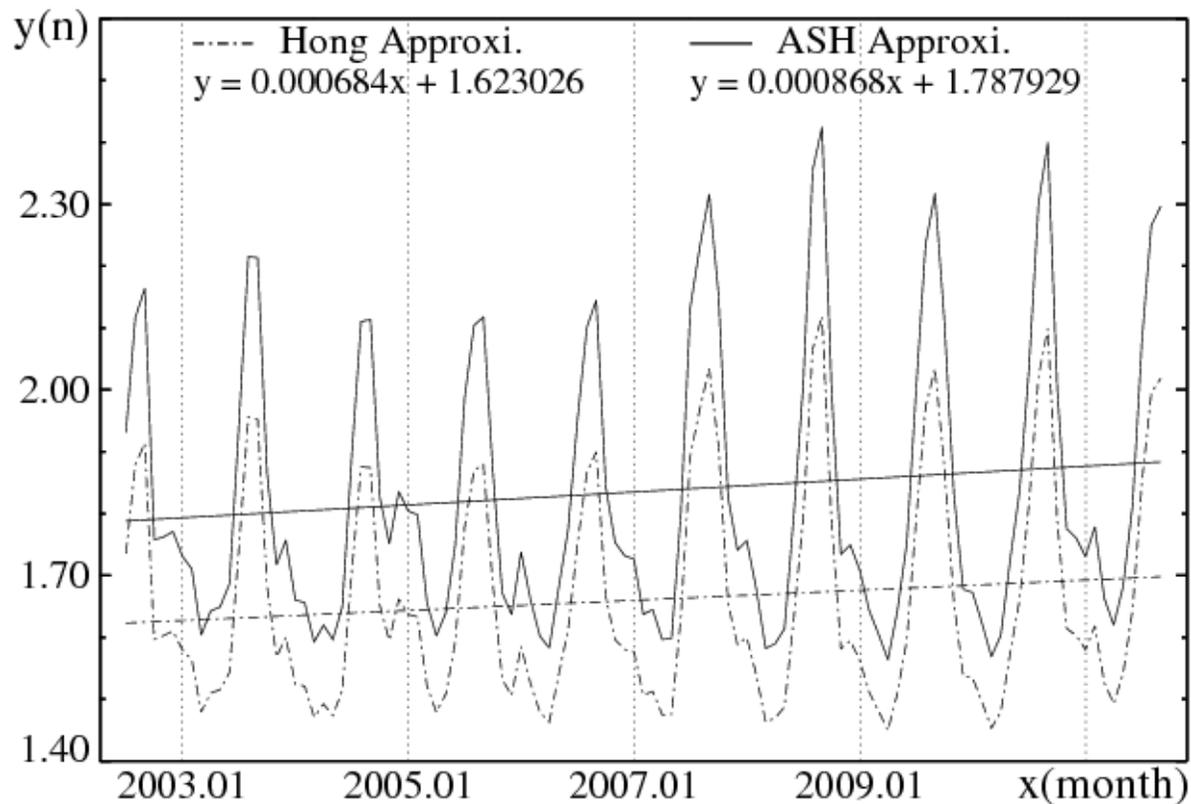


ASH approximation♪

Results (Refractive index)



- The time series of n from 2002 to 2011 show the increasing trend. The sea ice surface is melting because n is closer to that of water.





- The ASH and Hong approximations were applied to investigate the characteristics of the polar sea ice surface using the daily L3 25-km data of AMSR-E.
- The effective rough surface reflectivities were estimated for V and H polarizations.
- In turn, the surface roughness and refractive index were estimated using the ASH and Hong approximations.
- The ASH approximation produces a better result for sea ice than for water (Winter), while the Hong approximation is more effective for water than for sea ice in melting seasons (Summer).
- In conclusion, the combination of the ASH and Hong approximations is effective to estimate the small-scale roughness and refractive index for physical interpretations of sea ice melting due to climate change.
- This investigation supports the phenomenon of sea ice melting on the basis of the decreasing trend in the small-scale roughness and the increasing trend in the refractive index.

Thank you

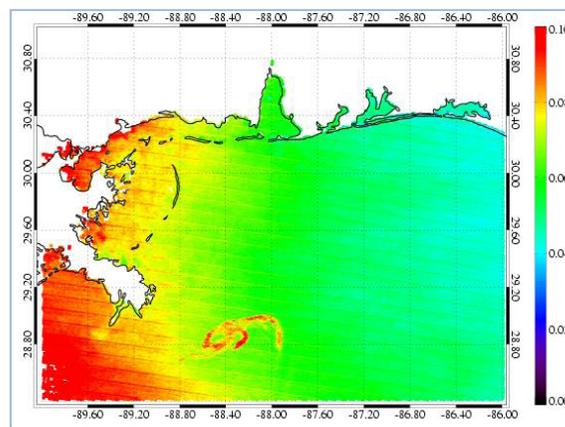




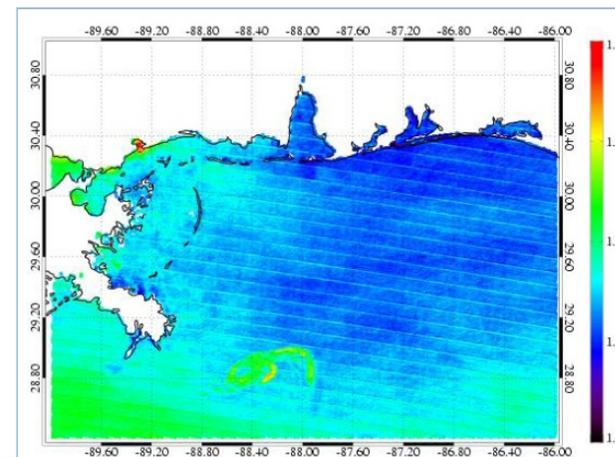
8. Oil spill detection



- Reflectivities
 - R_v over oil : 0.07 ~ 0.1, R_h : ~ 0.04

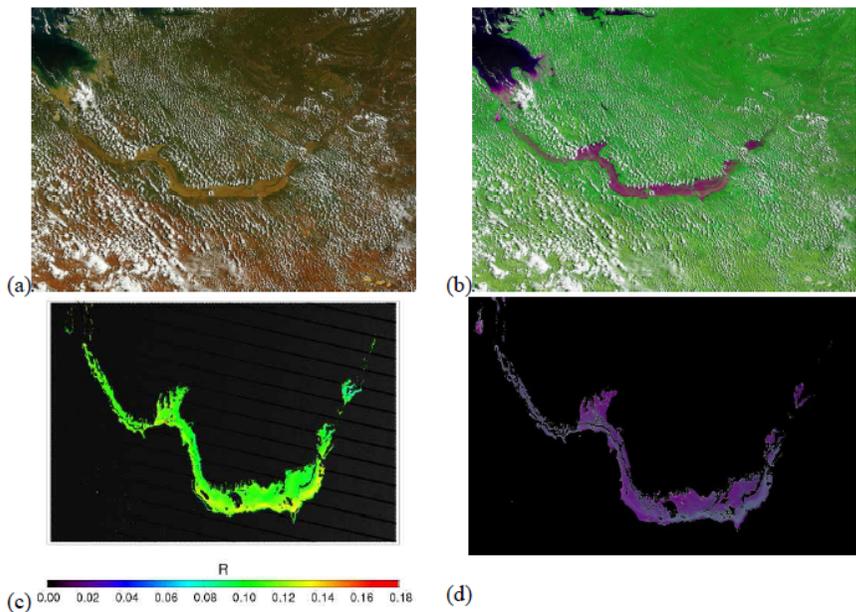


- Refractive index
 - n : 1.25 ~ 1.4 (oil), 1.1 ~ 1.25 (sea-water)
 - k : 0.45 ~ 0.5 (oil), 0.4 ~ 0.45 (water)
 - n values : physical, while k values are not physical but effective
 - **In review (2013)**

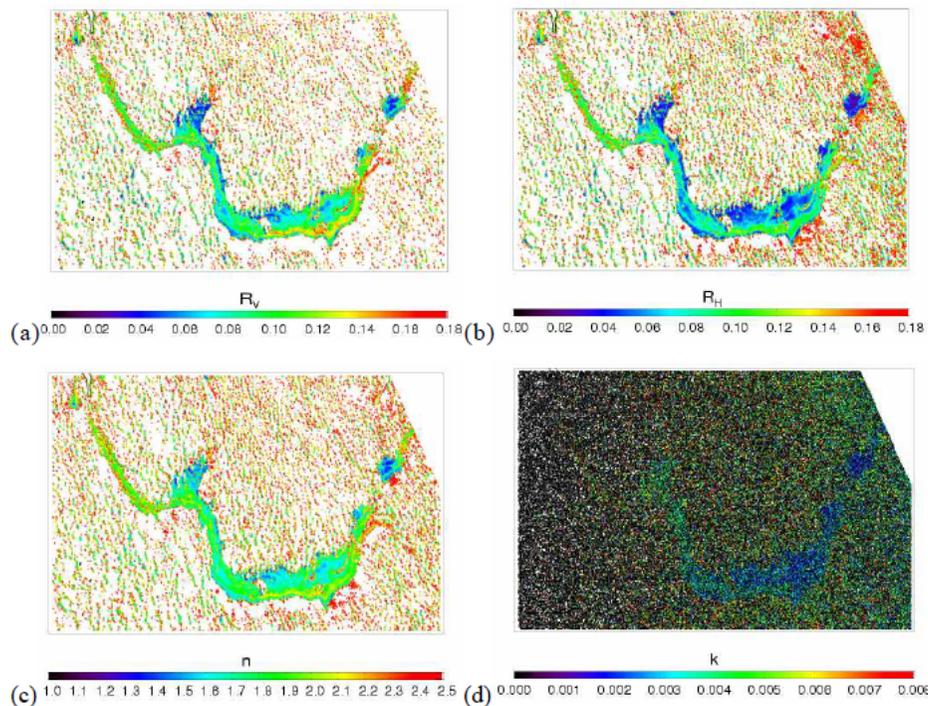




10. Flood monitoring

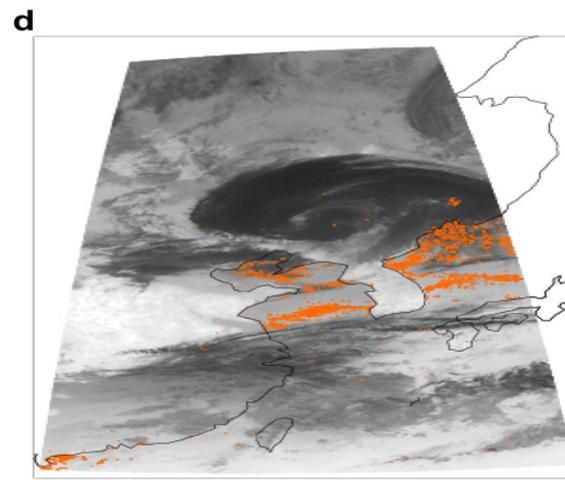
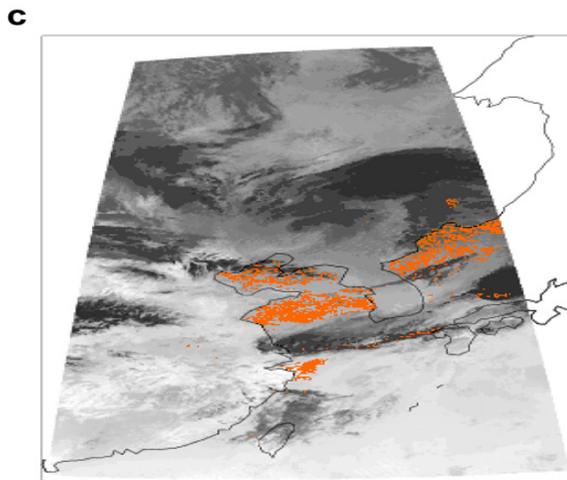
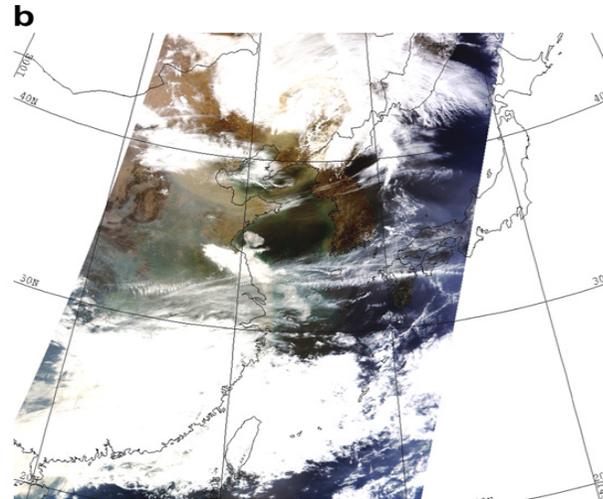
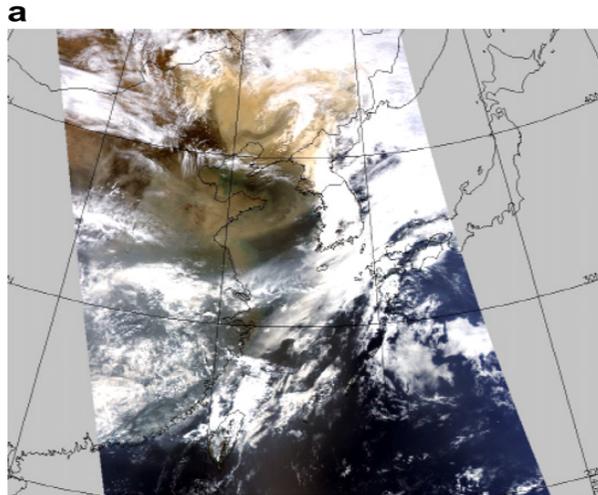


- In review
 - a flooding along the Fitzroy river area of Western Australia on March 17, 2011





4. Asian dust



- Hong, 2009 (Atmos. Environ.)



3. Emissivity model (Radiative transfer model)

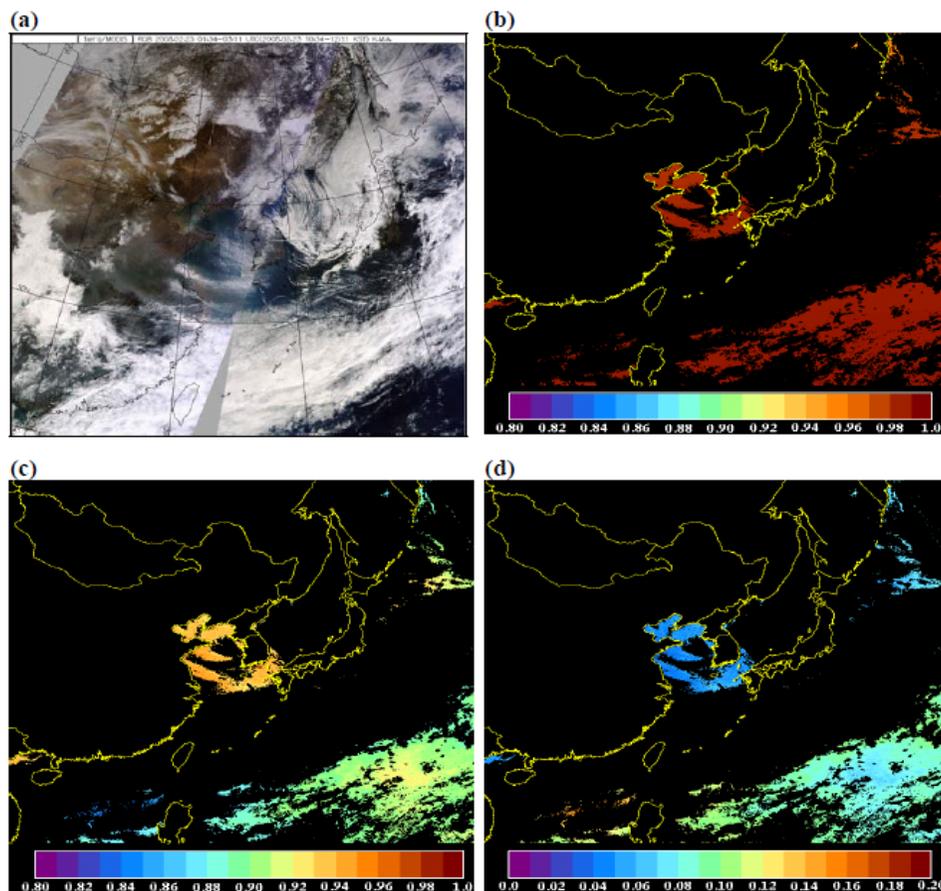


FIG. 3. MODIS RGB composite image and emissivities for the NW Pacific region. (a) MODIS

RGB image, (b) IE, (c) PE, and (d) Difference between IE and PE are illustrated, respectively.

The cloud is removed by cloud detection algorithm.

Hong et al. 2010 (J. Atmos. Ocean. Tech.)